

# The hydrostatic approximation for stratified fluids

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# Outline

- 1 Hydrostatic equations
- 2 Eddy diffusivity
- 3 Isopycnal coordinates
- 4 Results

## The initial set of equations

Motivated by geophysical flows, we consider the heterogeneous incompressible Euler equations with gravity force

$$(E) \quad \begin{cases} \partial_t \rho + \partial_x(\rho u) + \partial_z(\rho w) = 0, \\ \rho(\partial_t u + u\partial_x u + w\partial_z u) + \partial_x P = 0, \\ \rho(\partial_t w + u\partial_x w + w\partial_z w) + \partial_z P + g\rho = 0, \\ \partial_x u + \partial_z w = 0, \\ \text{boundary conditions.} \end{cases}$$

The velocity field is denoted  $(u, w) : \Omega \rightarrow \mathbb{R} \times \mathbb{R}$ , the density  $\rho : \Omega \rightarrow \mathbb{R}_*^+$ .

Reconstructing the pressure  $P : \Omega \rightarrow \mathbb{R}$  requires to solve an elliptic problem. Using instead the hydrostatic approximation, the system becomes ...

## The hydrostatic approximation

The heterogeneous, incompressible Euler equations with gravity force

$$(E) \quad \left\{ \begin{array}{l} \partial_t \rho + \partial_x(\rho u) + \partial_z(\rho w) = 0, \\ \rho(\partial_t u + u\partial_x u + w\partial_z u) + \partial_x P = 0, \\ \rho(\partial_t w + u\partial_x w + w\partial_z w) + \partial_z P + g\rho = 0, \\ \partial_x u + \partial_z w = 0, \\ \text{boundary conditions.} \end{array} \right.$$

Using the hydrostatic approximation, the system becomes

$$(H) \quad \left\{ \begin{array}{l} \partial_t \rho + \partial_x(\rho u) + \partial_z(\rho w) = 0, \\ \rho(\partial_t u + u\partial_x u + w\partial_z u) + \partial_x P = 0, \\ \partial_z P = -g\rho, \quad \iff \quad P = P|_{z=z_{\text{surf}}} + g \int_{\cdot}^{z_{\text{surf}}} \rho \, dz \\ \partial_z w = -\partial_x u, \quad \iff \quad w = w|_{z=z_{\text{bot}}} - \int_{z_{\text{bot}}}^{\cdot} \partial_x u \, dz \\ \text{boundary conditions.} \end{array} \right.$$

**Pb:** stability of shear flows  $(\rho, u)(t, x, z) = (\underline{\rho}(z), \underline{u}(z))?$

## Some stability results

### Homogeneous case: $\rho \equiv 1$ .

- Spectral stability of the linearized system about shear flows, under the Rayleigh criterion  $\underline{u}''(z) \neq 0$ . [Rayleigh (1880)]
- Lyapunov stability under the Rayleigh criterion. [Arnold '65]
- Well-posedness of the (nonlinear) system in Sobolev spaces under the Rayleigh criterion. [Grenier '99], [Brenier '03], [Masmoudi&Wong '12].

### Inhomogeneous case: $\partial_z \rho \neq 0$ .

- Spectral stability of the linearized system about shear flows, under the Miles and Howard criterion  $\frac{1}{4} |\underline{u}'(z)|^2 \leq \frac{-\rho'(z)}{\rho(z)}$ . [Miles '61][Howard '61]
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### Open problem

Well-posedness of the initial-value problem for the (linear or nonlinear) system?

## Possible source of instabilities

The hydrostatic limit is singular:

$$(H) \quad \left\{ \begin{array}{l} \partial_t \rho + u \partial_x \rho + w \partial_z \rho = 0, \\ \rho (\partial_t u + u \partial_x u + w \partial_z u) + \partial_x P = 0, \\ \partial_z P = -g \rho, \\ \partial_z w = -\partial_x u, \\ \text{boundary conditions.} \end{array} \right. \quad \begin{array}{l} \\ \\ \iff P = P|_{z=z_{\text{surf}}} + g \int_{z_{\text{bot}}}^{z_{\text{surf}}} \rho \, dz \\ \iff w = w|_{z=z_{\text{bot}}} - \int_{z_{\text{bot}}}^{\cdot} \partial_x u \, dz \end{array}$$

- Contributions from  $w \partial_z \rho$  and  $\partial_x P$  compensate in energy estimates if  $\rho > 0$ ,  $\partial_z \rho < 0$ . **Stable stratification helps.**
- There is no obvious way to deal with the contribution  $w \partial_z u$ . **Shear velocity hurts.**
- In the literature, theoretical results either
  - ① Use the analytic framework. e.g. [Kukavica, Temam, Vicol & Ziane '11]
  - ② Add regularization (viscosity, diffusivity). e.g. [Cao, Li & Titi '07+...]
  - ③ Relax the hydrostatic approximation [Desjardins, Lannes & Saut '21]

## Eddy viscosity and diffusivity

We add to the system the contributions of non-resolved eddies suggested by [Gent & McWilliams '90]. They act as additional “bolus” velocities.

$$(H_\kappa) \quad \begin{cases} \partial_t \rho + (u + u_\star) \partial_x \rho + (w + w_\star) \partial_z \rho = 0, \\ \rho (\partial_t u + (u + u_\star) \partial_x u + (w + w_\star) \partial_z u) + \partial_x P = 0, \\ \partial_z P = -g\rho, \quad \partial_z w = -\partial_x u, \\ \text{boundary conditions (free surface).} \end{cases}$$

with

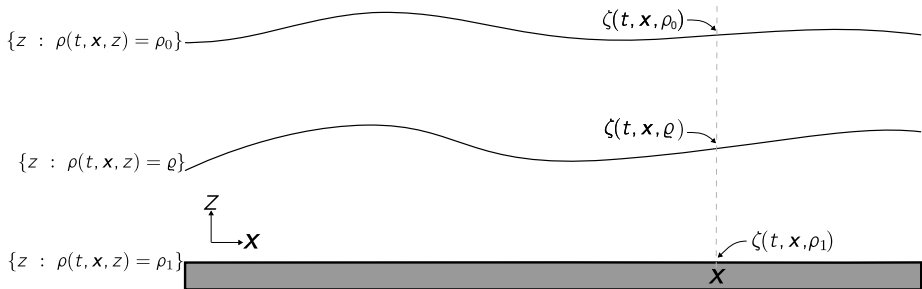
$$u_\star = \kappa \partial_z \left( \frac{\partial_x \rho}{\partial_z \rho} \right), \quad w_\star = -\kappa \partial_x \left( \frac{\partial_x \rho}{\partial_z \rho} \right),$$

where  $\kappa > 0$ .

# Isopycnal coordinates

We consider stratified flows and define the variable  $h(t, \mathbf{x}, \varrho) > 0$  through

$$h \stackrel{\text{def}}{=} -\partial_{\varrho}\zeta, \quad \rho(t, \mathbf{x}, \zeta(t, \mathbf{x}, \varrho)) = \varrho, \quad \zeta(t, \mathbf{x}, \rho(t, \mathbf{x}, z)) = z.$$





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The system

$$(H_{\kappa}) \quad \begin{cases} \partial_t \rho + (u + u_{\star}) \partial_x \rho + (w + w_{\star}) \partial_z \rho = 0, \\ \rho (\partial_t u + (u + u_{\star}) \partial_x u + (w + w_{\star}) \partial_z u) + \partial_x P = 0, \\ \partial_z P = -g\rho, \quad \partial_z w = -\partial_x u, \\ \text{boundary conditions (free surface).} \end{cases}$$

reads in isopycnal coordinates

$$(H_{\kappa}) \quad \begin{cases} \partial_t h + \partial_x (h(u + u_{\star})) = 0, \\ \partial_t u + (u + u_{\star}) \partial_x u + \partial_x (\mathcal{M}h) = 0, \end{cases}$$

where  $u_{\star} \stackrel{\text{def}}{=} \kappa \frac{-\partial_x h}{h}$  and

$$(\mathcal{M}h)(t, x, \varrho) \stackrel{\text{def}}{=} g \int_{\rho_0}^{\rho_1} \frac{\min(\varrho, \varrho')}{\varrho} h(t, x, \varrho') \, d\varrho'$$

# The system in isopycnal coordinates

The isopycnal formulation of the hydrostatic system is

$$(H_\kappa) \quad \begin{cases} \partial_t h + u \partial_x h + h \partial_x u = \kappa \partial_x^2 h, \\ \partial_t u + \left(u + \kappa \frac{-\partial_x h}{h}\right) \partial_x u + \partial_x (\mathcal{M}h) = 0, \end{cases}$$

where  $\mathcal{M}h(t, x, \varrho) \stackrel{\text{def}}{=} g \int_{\rho_0}^{\rho_1} \frac{\min(\varrho, \varrho')}{\varrho} h(t, x, \varrho') d\varrho'$ .

## Remarks:

- The parameterization of [Gent & McWilliams '90] is nice and simple.
- The advection in the variable  $z$  (or  $\varrho$ ) has disappeared.
- The domain is flattened.
- The system is easily discretized (multilayer framework) [Adim]
- The structure of the system is clarified:

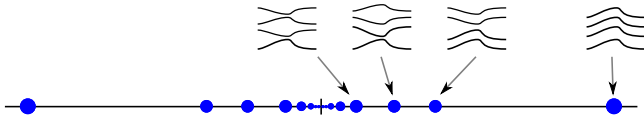
$$\partial_t \begin{pmatrix} h \\ u \end{pmatrix} + \begin{pmatrix} 0 & \text{Diag}(h) \\ \mathcal{M} & 0 \end{pmatrix} \partial_x \begin{pmatrix} h \\ u \end{pmatrix} + \begin{pmatrix} \text{Diag}(u) & 0 \\ 0 & \text{Diag}(u) \end{pmatrix} \partial_x \begin{pmatrix} h \\ u \end{pmatrix} = \kappa \begin{pmatrix} \partial_x^2 h \\ \frac{\partial_x h}{h} \end{pmatrix}$$

## Structure of the system

The isopycnal formulation of the hydrostatic system is

$$\partial_t \begin{pmatrix} h \\ u \end{pmatrix} + \begin{pmatrix} 0 & \text{Diag}(h) \\ \mathcal{M} & 0 \end{pmatrix} \partial_x \begin{pmatrix} h \\ u \end{pmatrix} + \begin{pmatrix} \text{Diag}(u) & 0 \\ 0 & \text{Diag}(u) \end{pmatrix} \partial_x \begin{pmatrix} h \\ u \end{pmatrix} = \kappa \begin{pmatrix} \partial_x^2 h \\ \partial_x^2 h \\ h \end{pmatrix}$$

- ① The first two terms generate wave propagation



- ② Advection contribution mixes modes when  $\partial_\rho u \neq 0$ .

- ③ The Gent&McWilliams contribution provides (partial) diffusion.<sup>1</sup>

<sup>1</sup>On a side note, if we define  $v = u - \kappa \frac{\partial_x h}{h}$ , then the system becomes

$$\partial_t \begin{pmatrix} h \\ v \end{pmatrix} + \begin{pmatrix} 0 & \text{Diag}(h) \\ \mathcal{M} & 0 \end{pmatrix} \partial_x \begin{pmatrix} h \\ v \end{pmatrix} + \begin{pmatrix} \text{Diag}(v) & 0 \\ 0 & \text{Diag}(v - \kappa \frac{\partial_x h}{h}) \end{pmatrix} \partial_x \begin{pmatrix} h \\ v \end{pmatrix} = \kappa \begin{pmatrix} 0 \\ \partial_x^2 v \end{pmatrix},$$

and the Gent&McWilliams contribution acts as an effective viscosity on the variable  $v$ .

## Some results

[R. Bianchini & VD]

For sufficiently regular data satisfying the stable stratification assumption

$$h|_{t=0} \geq h_0 > 0$$

and for any  $\kappa > 0$ , there exists a unique (classical) solution which we control on the time interval  $[0, T]$  with

$$T^{-1} = C \left( 1 + \kappa^{-1} (|\underline{u}'|_{L^2_\rho}^2 + M_0^2) \right),$$

where  $M_0$  is the size of the initial deviation from the shear flow equilibrium  $(\underline{\rho}(\varrho), \underline{u}(\varrho))$ , and  $C$  depends only on  $M_0$ ,  $h_0$  and the size of  $(\underline{\rho}(\varrho), \underline{u}(\varrho))$ .

[R. Bianchini & VD]

As long as this solution is controlled, we have strong convergence of the corresponding solutions to the incompressible Euler equations towards this solution in the limit of shallow water aspect ratio  $L_z/L_x \ll 1$ .

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## Thoughts to go

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Thank you for your attention